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37 Introduction and Objectives of the Technical Report

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39 Land changes in the coastal zone and high sea-level rise rates are exposing lowland areas to more

40 frequent events of saltwater intrusion, flooding, and rapid shoreline erosion, magnifying the

41 negative effects of coastal storms and storm surge. Louisiana is particularly sensitive to sea-

42 level rise due to the unique geology of the State's Delta and Chenier Plains. There is a pressing

- 43 need to integrate up-to-date sea-level rise data into planning and engineering activities to
- 44 anticipate coastal land loss patterns, protect coastal communities and adequately design
- 45 restoration projects.
- 46

47 Projections for future sea-level rise and concurrent coastal vulnerability estimates are numerous
48 and variable. As sea-level rise projections are refined and confidence in future sea level rise
49 estimates increase, policy development in state and regional planning efforts will become more

- 50 mature. Until a more precise picture of sea-level rise trends emerges, adaptability and flexibility
- 51 will be key components to ensure the incorporation of the most current and accurate sea-level
- 52 rise projections into project planning and policy development for the Louisiana coast.
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- 54 The objective of the technical report is to make recommendations for incorporating sea-level rise
- 55 into the planning and engineering of habitat restoration and storm protection projects.
- 56 Specifically the technical report is structured to:
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- Deductively summarize the state of the science on the patterns of increase in the surface of the global ocean, regional Gulf of Mexico and local coastal waters to support a recommendation of the rate(s) of anticipated sea-level rise most appropriate for incorporating into project analysis and design, and
- 61 62 63
- Describe how recommended rate(s) of local sea-level rise should be combined with the highly variable spatial patterns in coastal subsidence and wetland vertical accretion to predict relative sea-level rise at specific points in the Louisiana coastal zone.
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67 This summary provides an overview of the findings of the technical report and the

recommendations that result from those findings. Both the summary and full technical report

69 should be considered as 'living documents.' As global, regional and local estimates of sea-level

rise are constantly changing as new data become available, CPRA Louisiana Applied Coastal

- 71 Engineering & Science (LACES) Division plans to update this report as necessary.
- 72

73 Background

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75 Sea-level rise is caused by a variety of dynamic interactions, and is influenced by atmospheric,

76 geologic, oceanic, and astronomical changes, whether natural or anthropogenic. Estimates of

77 past and projections of future SLR are dependent on the interplay between these factors.

78 Challenges arise when attempts are made to identify changes in sea level that occur over varying

79 time scales; therefore, it is important to isolate the long term historical trends from the

80 background of regular natural cycles to make confident predictions about future trends.

81

82 Global sea level rise (GSLR) is a combination of thermal expansion, or the steric component,

83 and the freshwater influx, or the eustatic component. To accurately determine the long-term

84 historical trend of global sea level rise, researchers must first identify and remove the effects of

85 natural forcings on the sea level observation data to evaluate the steric and eustatic components

86 only. Natural forcings are generally cyclical in nature and can be caused by atmospheric,

87 meteorological or alterations in the rotation of the earth and moon. Because these patterns are

88 reflected in monitoring data and can mask long term trends, it is preferable to have data records 89

long enough to encompass several years so that these patterns can be identified and removed

90 from the analysis.

91

92 While it is important for coastal managers to consider GSLR, local sea level rise or relative sea

93 level rise (RSLR) is more relevant in coastal planning. In any coastal zone, the actual rate of

94 SLR is a combination of GSLR and local processes including natural cycles, glacial isostatic

95 adjustment (GIA), subsidence, accretion and erosion of shorelines and coastal marshes. These

96 influences result in a RSLR that may be quite different from GSLR. Therefore, it is necessary to 97

evaluate GSLR trends and then to focus on local conditions in the Gulf of Mexico offshore of 98

southern Louisiana to inform recommendations on estimating and incorporating local RSLR

- 99 trends into project planning and design.
- 100

101 Sea Level Rise Measurement

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103 Oceanographic data are gathered from three primary sources to determine the global sea-level

104 rise rate: tide gauges, *in situ* measuring devices and satellite altimetry readings. In the United

105 States, the National Oceanographic and Atmospheric Administration operate 128 long-term

106 National Water Level Observation Network tide gauge stations that monitor monthly mean sea-

107 level data. Measurements from this network are used in conjunction with other tide gauges

108 around the globe to calculate the rate of global sea-level rise.

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110 The distribution of the *in situ* devices, such as the Argo float array and the National

Oceanographic and Atmospheric Administration's bathythermograph monitoring fleet are used 111

112 to define the specific contributions of the eustatic and steric components of current sea-level rise

measurements. If distribution of the *in situ* devices is sparse in certain water bodies, such the 113

114 Gulf of Mexico, it becomes more difficult to establish the relative contribution of eustatic and

115 steric components of current sea-level rise measurements and complicate predictions of future

116 sea-level rise. Thus, the Argo float data is of limited use to Louisiana's coastal zone.

118 Satellite altimetry measurements can also be used to measure sea level change. A series of

satellite missions beginning with TOPEX/Poseidon in 1992 and continuing today with Jason 1

120 and 2 provide estimates of global mean sea level rise. It should be noted, however, that satellite

- altimetry data must be processed and corrected to account for influences such as GIA and
- seasonality. Additionally, the relative short period of record for satellite altimetry data makes it

123 difficult to discern long term trends in sea level change.

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125 Historical Sea-Level Rise Rates

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Estimating a historical rate of global sea-level rise is entirely dependent on the specific time period and measurement method. Much of the debate about calculating historical sea-level rise trends, including rates and acceleration, results from which stations are chosen and over which specific time period they are analyzed.

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132 Most efforts to calculate an average rate of global sea-level rise for the 20th century have

133 concluded a value of less than 2 mm/yr (0.08 inches/yr; e.g. Church et al. 2004; Church and

134 White 2006). In comparison, NOAA's Laboratory for Satellite Altimetry has calculated an

estimated overall 1993-2010 global sea-level rise rate of 2.8 ± 0.4 mm/yr (data accessed on 9

136 September 2011). It is important to note that these calculations assume an equal and constant

137 sea-level rise trend. However, examination of the historical and a large body of recent literature

point to a rate of global sea-level rise since the late nineteenth century that has been accelerating.

139

140 To illustrate this point, Figure S1 shows the data set by which Church and White (2006)

141 calculated the frequently-cited 1.7 ± 0.3 mm/yr (0.067 ± 0.012 inches/yr) average global sea-

142 level rise for the twentieth century. That conclusion was down-revised to 1.5 mm/yr (0.059

143 inches/yr) for the whole 1860-2009 time period in Church and White (2011). In those two

- papers Church and White discussed that in the least any linear analysis needed to recognize
- several visually-obvious inflection points in the data in the mid-1930s, the 1960s and the 1980s,
- and that rates calculated for the intervening years between those points might differ significantly.
- However, in both their 2006 and 2011 papers Church and White also defined a non-linear,
- accelerating fit to the full data set. Acceleration in the historical data was also documented by
 Jevrejeva et al. (2008) and Woodworth et al. (2009). Differences in forecasting a linear vs.
- 149 Jevrejeva et al. (2008) and woodworth et al. (2009). Differences in forecasting a linear vs. 150 accelerating function into the future can become significant and the use of one prediction over
- another should be carefully considered.
- 152

153 Satellite altimetry data specific to the Gulf of Mexico do show a slightly lower rate of sea surface

154 elevation change than the global average over the past twenty years (Figure S2). To improve

155 future estimates of future Gulf-wide sea-level rise, we are exploring the reasons for this

156 observation. Satellite altimetry-based estimates of sea-level rise in the Gulf of Mexico are also

157 highly variable at more local levels, with the general pattern of higher rates of sea-level rise in

158 the center of the Gulf and lower rates in both the eastern and western margins.

159

160 At a more local level, there appears to be a strong east-west gradient in derived 1992-2010 Gulf

161 of Mexico sea-level rise values for the near-shore coastal waters in southern Louisiana (Figure

- 162 S3), with values ranging from a high southeast of the Balize Delta and consistently decreasing
- 163 both westward to a low south of the mouth of the Sabine River and northward into the Lake



Figure S1. Sea-level rise data from Church and White (2011) for the 1860-2009 time period was used to calculate an 1880-2009 linear trend of 1.5 mm/yr (0.059 inches/yr) and the non-linear, accelerating trend discussed in the text of the Technical Report.

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Borgne and Lake Pontchartrain systems. The range of values is significant, with the highestderived sea-level rise rates being 58% higher than the lowest values.

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173 **Projections of Future Sea-Level Rise**

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The largest uncertainty in predicting eustatic sea-level rise over the next century is how ice sheets will respond to changes in temperature (Allison et al. 2009. Since satellite altimetry and

- sheets will respond to changes in temperature (Amson et al. 2009. Since saterine attimetry and any angle to more clearly investigate the
- 177 gravimetric data have become available, researchers are able to more clearly investigate the
- dynamics of ice sheets. Current estimates for global sea-level rise by 2100 range from 0.5 to 2
- 179 meters (1.6' 6.6'; Grinsted et al. 2009; Pfeffer et al. 2008; Rahmstorf 2007; Vermeer &
- 180 Rahmstorf 2009). Much weight recently has been placed on the Rahmstorf (2007) predictions of
- 181 a range in global sea- level rise of 0.5-1.4 meters (1.6' 4.6') by 2100, with 1 meter (3.3') being
- 182 the most likely. With respect to sea level rise rate acceleration, we have chosen to follow the
- 183 weight of scientific opinion that sea-level rise is in fact accelerating. However, LACES will 184 follow this issue closely in the future and will revision our recommendations accordingly.
- 185

186 In comparison to the projections of global sea-level rise, there has been very little work done to

- 187 predict the specific change in the Gulf of Mexico water surface for the rest of this century. Until
- these regional investigations are performed, anticipated sea-level changes in the Gulf of Mexico



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Figure S2. Gulf of Mexico sea level values sourced in satellite altimetry data from the National
Oceanographic and Atmospheric Administration's Laboratory for Satellite Altimetry, accessed on 11
March 2011, illustrate an overall sea-level rise rate for the Gulf of Mexico lower than the global sea-level
rise trend calculated for the same time period (2.8 ± 0.4 mm/yr).

196 197

198 must be primarily extrapolated from satellite altimetry or tide gauges, which can be less reliable 199 due to the limited period of record. Modeling of global climate change scenarios suggests that 200 the Gulf of Mexico will respond similarly to the global ocean (Meehl et al. 2007), and it appears 201 that projections of global sea-level rise are appropriate to carry into the region.

202

In order to calculate a predicted future sea-surface elevation offshore of coastal Louisiana, and based on the literature review above, the Technical Report recommends that CPRA staff assume that Gulf sea-level rise will be 1 meter (3.3') by 2100, with a bounding range of 0.5 - 1.5 meters (1.6' - 4.9'). While this recommendation results from the independent assessment of the available data presented in the Technical Report, it is also consistent with similar efforts ongoing in other states. This recommendation is only part of the overall prediction of future relative sealevel rise, and must be combined with predictions of subsidence and marsh vertical accretion.

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1.86	2.06	2.35	2.57	2.72	2.81	2.83	2.93	3.1 ●	3.29	3.5 O	3.67	3.81	4.04	4.35 <mark>0</mark>	4.47	4.42	4.11	3.55
1.82	2.07	2.47	2.77	2.95	3.05	3.08	3.19	3.36	3.55	3.74	3.94 <mark>O</mark>	4.14	4.44	4.82	4.99	4.94	4.55	3.82
1.67	1.95	2.46	2.86	3.14	3.32	3.38	3.47	3.59	3.73	3.9 <mark>O</mark>	4.14	4.44	4.81 	5.24	5.41 ●	5.31	4.83	3.95
1.41	1.71	2.32	2.85	3.31 ●	3.59	3.71	3.77	3.78	3.85	3.99	4.27	4.71	5.16	5.62	5.75	5.55	4.95 ●	3.96

Figure S3. Sea-level rise rates for discrete points offshore of southern Louisiana show significant eastwest variation, with values highest offshore of the Balize Delta and trending lower moving west across the front of the Chenier Plain. Data were derived by USGS from satellite altimetry data for center points of the analysis grid shown in Figure 12 of the Technical Report, covering the 1992.96-2010.01 period of record. Figure from the Project Effects Modeling Team addressing the 2012 revision of the Louisiana 218 Integrated Master Plan.

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221 Having discussed global sea-level rise projections, any effort to confidently incorporate potential 222 sea-level rise impacts on coastal wetlands into planning must account for the sum of factors 223 influencing relative sea-level rise: 1) the change in the surface elevation of the Gulf of Mexico, 224 which is the primary topic of this document; 2) local land surface elevation change, which in 225 Louisiana is exclusively represented as subsidence; and 3) marsh vertical accretion, which can 226 offset some sea-level rise impacts.

227

228 Subsidence is a significant driver of relative sea-level rise in southern Louisiana, and is most

229 likely the principal driver in southeast Louisiana for the near-term. While rates of subsidence are

- 230 highly variable across the Louisiana coastal zone (Figure S4), our understanding of the exact
- 231 rates of subsidence at the local level is very limited. Ongoing work by the Louisiana Geological
- 232 Survey, commissioned by LACES, will summarize our understanding of the geological
- 233 framework underlying south Louisiana as well as provide an overview of historical rates of 234 subsidence across the landscape. This work will further build on the summary report
- 235 Understanding Subsidence in Coastal Louisiana (Reed and Yuill 2009), prepared for the State
- 236 and the US Army Corps of Engineers by the Louisiana Coastal Area Science and Technology
- Program Office. This information, together with monitoring data from the Coastal Wetland 237
- Planning, Protection and Restoration Act Program's Coastal Reference Monitoring System -238
- 239 *Wetlands* stations, will help to determine the predicted ranges of subsidence.

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Figure S4. Map of projected subsidence ranges for south Louisiana generated by the Subsidence
 Advisory Panel for the Louisiana CPRA Master Plan 2012 Update, following a meeting on 14 October
 2010.

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248 Our understanding of marsh vertical accretion is likewise evolving. The ability of marshes to 249 keep up with moderate levels of sea-level rise via accretion of both mineral and organic soil 250 material has long been understood (see summary in Mitsch and Gosselink 2000). Typically coastal marshes have a range of "optimum depth" and, within this range, the marshes will 251 252 respond positively to a sea-level rise. Once they reach an estuary specific rate of rise where 253 accretion is no longer possible, the marshes will then drown. While this science is nascent at 254 present, it promises to be a significant contribution to predicting local net relative sea-level rise 255 in Louisiana's coastal wetlands.

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This paper has focused on global and regional sea level rise. However, the natural processes of accretion and subsidence are critically important to how the coast will evolve over time. As the science continues to evolve and we begin to get a better understanding of how the processes of accretion and subsidence will affect our coastal zone, future research and synthesis papers will be prepared to address these two issues.

263 Summary and Recommendations

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265 The scientific literature indicates that the rate of global SLR (GSLR) has been increasing steadily

266 over the past several centuries. This may be seen in an increase from a 20th Century linear

267 average based on tide gauge data of $1.7 \pm 0.4 \text{ mm/yr.}(0.070 \pm 0.016 \text{ inches/yr.})$ to a linear 268 estimate for the past eighteen years of $2.9 \pm 0.4 \text{ mm/yr.}(0.11 \pm 0.016 \text{ inches/yr.})$ in the Gulf of

269 Mexico based on satellite altimetry. Although direct comparison of the two techniques supports

the validity of the altimeter readings, there is some concern regarding the short period of record

271 for the altimetry data. However, evidence suggests that SLR for the available period of record is

best represented as a single, non-linear function, which has important implications for relating

RSLR and GSLR estimates, and especially for assumptions of the differential representing localland surface change.

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276 More important for CPRA planning purposes is the projection of future GSLR. Based on the

available data, LACES recommends that any SLR modeling scenarios models for state

restoration projects assume a 1-meter (3.3') MSL rise by 2100 compared to the late 1980s and

should be bracketed by GSLR ranges of 0.5-1.5 meters (1.4'-4.9') by 2100. The specific

recommendation for factoring in the range of GSLR into local calculations of RSLR is givenbelow.

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Recommendations for Calculating RSLR in Coastal Louisiana 284

Based on the information presented to this point, it is our recommendation that when
participating in project planning and design activities, local RSLR be calculated using the
following procedure to populate the variables of the generalized RSLR equation

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 $E(t) = a^{*}t + b^{*}t^{2} + S.$ (Eqn. 7)

1. Use local observations of historical sea-level rise from contemporary satellite altimetry (Figure 16) just offshore of coastal Louisiana, in order to account for the substantial east-west gradient in documented rates. Specifically, we recommend using an average of the three most proximate points shown in Figure S3. This is the rate of SLR (mm/yr.) and variable (a) from the generalized equation.

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2. Calculate the acceleration constant that assumes a MSL increase of 1 meter (3.3') by
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2100 as the most heavily-weighted project alternative, while also testing MSL increases
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of 0.5 meters (1.6') and 1.5 meters (4.9') to account for uncertainty in the literature.
300

This provides the change in water levels over time at a project location. To localize further, 302

3. Add in local subsidence values obtained from the most proximate local source, which is variable (S)in Equation 7.

In order to predict the persistence the coastal wetland, and specifically the persistence of the
wetland surface or conversely marsh surface collapse and drowning, a fourth step is necessary.

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4. Use the sum of #s 1-3 above to establish an inundation function, especially the rate of
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4. Use the sum of #s 1-3 above to establish an inundation function, especially the rate of
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- 315 As discussed in Section 2.4.1., predicting future RSLR must account for the acceleration
- 316 constants (variable **b**) being specific to NRC (1987) acceleration scenarios having a starting
- 317 point of 1986. Appendix D of the report shows specifically how the variables discussed feed
- into a refined version of Equation 8 that accounts for that specific starting point.
- 319

320 References

- 321
- Allison, I., R. B. Alley, H. A. Fricker, R. H. Thomas and R. C. Warner. (2009). Review: Ice
 sheet mass balance and sea level. *Antarctic Science* 21: 413-426.
- Church, J. A. and N. J. White. (2006). A 20th century acceleration in global sea-level rise.
 Geophysical Research Letters 33: L01602, doi:10.1029/2005GL024826.
- Church, J.A., and N.J. White. (2011). Sea-level rise from the late 19th to the early 21st century.
 Survey of Geophysics online publication, 30 March 2011.
- Church, J. A., N. J. White, R. Coleman, K. Lambeck and J. X. Mitrovica. (2004). Estimates of
 the regional distribution of sea-level rise over the 1950-2000 period. *Journal of Climate*, 17,
 pp 2609-2625.
- Grinsted, A., Moore, J. C., and Jevrejeva, S. (2009). Reconstructing sea level from paleo and
 projected temperatures 200 to 2100 ad. *Climate Dynamics*, 34(4), 461-472.
- Houston, J.R. and R.G. Dean (2011). Sea-level acceleration based on U.S. Tide Gauges and
 Extensions of Previous Global-Gauge Analyses. *Journal of Coastal Research* 27: 409-417.
- Jevrejeva, S. A., J. C. Moore and A. Grinsted. (2008). Relative importance of mass and volume
 changes to global sea-level rise. Journal of Geophysical Research, 113, D08105,
 doi:10.1029/2007JD009208.
- Meehl, G.A., T.F. Stocker, W.D. Collins, P. Friedlingstein, A.T. Gaye, J.M. Gregory, A. Kitoh,
 R. Knutti, J.M. Murphy, A. Noda, S.C.B. Raper, I.G. Watterson, A.J. Weaver and Z.-C.
- Zhao, 2007: Global Climate Projections. In: *Climate Change 2007: The Physical Science Basis. Contribution of Working Group I to the Fourth Assessment Report of the*
- *Intergovernmental Panel on Climate Change* [Solomon, S., D. Qin, M. Manning, Z. Chen,
 M. Marquis, K.B. Averyt, M. Tignor and H.L. Miller (eds.)]. Cambridge University Press,
- M. Marquis, K.B. Averyt, M. Tignor and H.L. Miller (eds.)]. Cambridge University Press,
 Cambridge, United Kingdom and New York, NY, USA.
- Mitsch, W. J. and J. G. Gosselink. (2000). Wetlands. John Wiley & Sons, Inc. New York, NY.
 920pp.
- National Research Council. (1987). Responding to Changes in Sea Level: Engineering
 Implications. National Academies Press, Washington, D.C. 160 pp. Pfeffer, W.T., J. T.
 Harper and S. O. Neel. (2008). Kinematic constraints on glacier contributions to 21st-century
 sea-level rise. Science, Vol 321, pp 1340-1343.
- Rahmstorf, S. (2007). A semi-empirical approach to projecting future sea-level rise. Science, Vol 315, pp 368-370.
- Reed, D. J., and B. Yuill. (2009). Understanding subsidence in Coastal Louisiana. Ponchartrain
 Institute for Environmental Sciences, University of New Orleans, New Orleans, LA. 69 pp.
- United States Army Corps of Engineers. (2009). Water resource policies and authorities
 incorporating sea-level change considerations in civil works programs. Circular No. 1165-2 211.
- Vermeer, M. S., and S. Rahmstorf. (2009). Global sea level linked to global temperature.
 Proceedings of the National Academy of Sciences 106: 21527-21532.
- Woodworth, P.L., N.J. White, S. Jevrejeva, S.J. Holgate, J.A. Church and W.R. Gehrels. 2009.
 Evidence for the accelerations of sea level on multi-decade and century timescales.
- 362 International *Journal of Climatology* **29**: 777-789.
- 363