

Climate change impacts in Latin America and the Caribbean and their implications for development

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Abstract This paper synthesizes what is known about the physical and biophysical impacts of climate change and their consequences for societies and development under different levels of global warming in Latin America and the Caribbean (LAC). Projections show increasing mean temperatures by up to 4.5 °C compared to pre-industrial by the end of this century across LAC. Associated physical impacts include altered precipitation regimes, a strong increase in heat extremes, higher risks of droughts and increasing aridity. Moreover, the mean intensity of tropical cyclones, as well as the frequency

of the most intense storms, is projected to increase while sea levels are expected to rise by ~0.2–1.1 mm depending on warming level and region. Tropical glacier volume is found to decrease substantially, with almost complete deglaciation under high warming levels. The much larger glaciers in the southern Andes are less sensitive to warming and shrink on slower timescales. Runoff is projected to be reduced in Central America, the southern Amazon basin and southernmost South America, while river discharge may increase in the western Amazon basin and in the Andes in the wet season. However, in many regions, there is uncertainty in the direction of these changes as a result of uncertain precipitation projections and differences in hydrological models. Climate change will also reduce agricultural yields, livestock and fisheries, although there may be opportunities such as increasing rice yield in several LAC countries or higher fish catch potential in the southernmost South American waters. Species range shifts threaten terrestrial biodiversity, and there is a substantial risk of Amazon rainforest degradation with continuing warming. Coral reefs are at increasing risk of annual bleaching events from 2040 to 2050 onwards irrespective of the climate scenario. These physical and biophysical climate change impacts challenge human livelihoods through, e.g., decreasing income from fisheries, agriculture or tourism. Furthermore, there is evidence that human health, coastal infrastructures and energy systems are also negatively affected. This paper concludes that LAC will be severely affected by climate change, even under lower levels of warming, due to the potential for impacts to occur simultaneously and compound one another.

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Introduction

Despite increasing numbers of climate change impact studies in Latin America and the Caribbean (LAC, defined here as the region encompassing the South American continent, Central America, the Caribbean islands and Mexico) summarized in the fifth assessment report of the IPCC (chapters by Magrin et al. 2014; Romero-Lankao et al. 2014; Nurse et al. 2014), knowledge on how climate change affects different subregions and sectors is fragmented. The objective of this paper is to analyze physical and biophysical impacts of climate change in LAC, and their consequences for societies and development, in an integrated way, thereby updating and extending the analysis presented in Schellnhuber et al. (2014). This paper adds to the recent work by the IPCC: Firstly, we consider the latest science published after the IPCC cutoff deadlines. Secondly, we consider climate impacts through a development lense and therefore devote more depth to sectors, subregions and systems which seem most relevant from that risk perspective. Thirdly, we synthesize climate impacts as a function of different levels of global warming at the end of the twenty-first century compared to pre-industrial levels (1.5, 2, 3 and 4 °C, cf. Table ESM.1 and Schellnhuber et al. 2014: Appendix 4) to show what these global warming levels mean for LAC. We focus on warming levels in general and on a 2 and 4 °C world (basically following the warming pathways of the scenarios RCP2.6 and RCP8.5, respectively) in particular since these are important elements of global climate change negotiations, but in such global discourses, regional implications of a global warming level often remain fuzzy.

We combine original data analyses, model projections and meta-analyses of published studies with a comprehensive literature review (the methodological approach is presented in Schellnhuber et al. 2014, especially in Appendices A.1–3). The impacts of climate change in LAC under different warming levels are synthesized in Fig. 5 and Table ESM.1. The temperature, precipitation, evapotranspiration and aridity projections are based on five CMIP5 GCMs (as selected by Warszawski et al. (2014), see Schellnhuber et al. 2014: Appendix A.1). We focus exclusively on climate impacts and do not consider possibilities of adaptation options decreasing these impacts unless otherwise stated.

Social, economic and demographic profile of the Latin America and Caribbean region and vulnerabilities to climate change

The countries in the LAC region differ greatly in their economic and demographic profile (see Table ESM.2). The

region comprises a population of 588 million (in 2013) which is expected to rise to 660 million by 2025 (World Bank 2014a). The region is highly urbanized, and in 2010, the urban population accounted for 78.8 % of the total and is expected to increase further (ECLAC 2014). The current GDP of the region is estimated at US\$ 5.467 trillion (in 2012) (World Bank 2014b) with a GNI per capita of US\$ 9536 in 2013 (World Bank 2014c). In 2012, approximately 25.03 % of the population was living in poverty and 12.02 % in extreme poverty or deprivation (World Bank 2014d). In 2010, the rural poverty rate was twice as high as that of urban areas; when considering extreme poverty, it was four times as high (IFAD 2013). Close to 60 % of the population in extreme poverty lives in rural areas (RIMISP 2011). Moreover, ethnicity, gender and age correlate with poverty. In seven countries for which data are available, the poverty rate is 1.2–3.4 times higher for indigenous and afrodescendent groups than for the rest of the population and poverty rates are 1.7 times higher among minors under 15 than in adults and 1.15 times greater among women than men (ECLAC and UNFPA 2009).

Climate change is expected to accentuate preexisting vulnerabilities in LAC. Several million people live in the path of hurricanes and in low-elevation coastal zones rendering them vulnerable to sea-level rise, storm surges and coastal flooding (McGranahan et al. 2007; Trab Nielsen 2010). LAC's 64,000-km coastline is one of the most densely populated in the world (Sale et al. 2008), and several countries have a large share of their urban population living in areas elevated less than five meters above sea level (CIESIN 2011; cf. Table ESM.3). Moreover, people living in slums built on steep slopes and with poor drainage systems (Douglas et al. 2008), and certain population groups [such as short-term or chronically poor people (Ahmed et al. 2009; Hardoy and Pandiella 2009; Hertel et al. 2010) and women-headed households or children (Kumar and Quisumbing 2011)] are particularly exposed to climate change risks. The rural poor in general, and indigenous groups in particular, are especially vulnerable to climate change because of their reliance on small-scale, rain-fed agriculture, natural resources, traditional knowledge systems and culture (Kronik and Verner 2010; Hoffman and Grigera 2013) and their poor access to infrastructure and technology (Feldt 2011). Many of these population groups also have limited political influence, fewer capabilities and opportunities for participating in decision and policy making and are less able to leverage government support to adapt to climate change (Hardoy and Pandiella 2009; Moser et al. 2010). Climate change may affect also the credibility of elders and traditional leaders, as their authority to predict the natural seasonality is challenged (Kronik and Verner 2010). Finally, a high proportion of the urban population in LAC lives in a few

very large cities. National economies, employment patterns and government capacities are also strongly dependent on these large cities making them extremely vulnerable to climate-related disasters (Hardoy and Pandiella 2009). In the following, we revise the state of the art of the climate component of the vulnerabilities stated here.

Regional patterns of climate change

Projected temperature changes

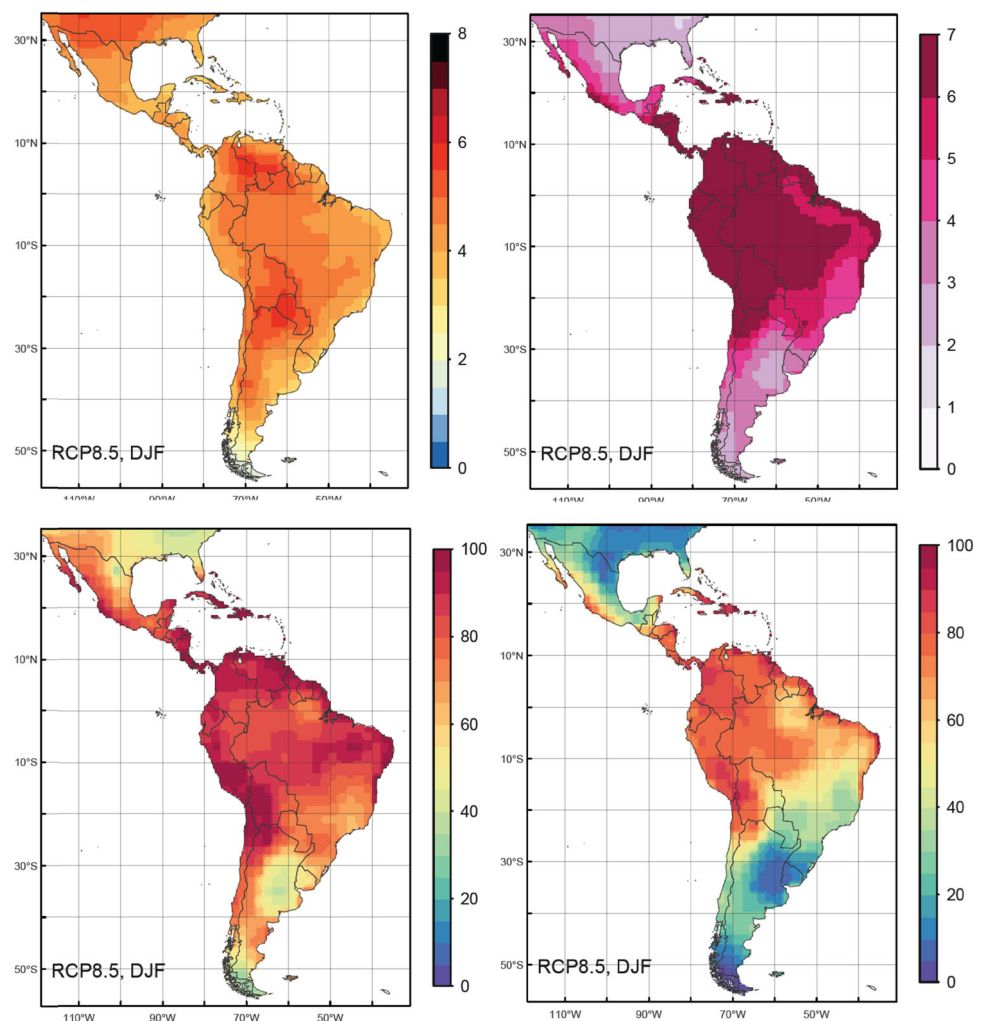
The multi-model mean austral warming for 2071–2099 is about ~ 1.5 °C in a 2 °C world and ~ 4.5 °C in a 4 °C world compared to 1951–1980 (Fig. ESM.1) and shows a rather uniform pattern, with more warming toward the interior of the continent (Fig. 1 top left panel, Fig. ESM.2). The normalized warming [i.e., the warming expressed in terms of the local year-to-year natural variability (cf.

Schellnhuber et al. 2014: Appendix A.1)—top right panel of Fig. 1, Fig. ESM.2] indicates how the projected warming compares to the natural fluctuations a particular region has experienced during the period 1951–1980 (Hansen et al. 2012; Coumou and Robinson 2013; Mora et al. 2013a). The tropics will see the strongest increase in normalized monthly summer temperatures, since historic year-to-year fluctuations are relatively small, indicating a new climatic regime for the tropical parts of LAC (top right panel of Fig. 1, Fig. ESM.2). Subtropical regions in the south (northern Argentina) and the north (Mexico) are expected to see a much less pronounced shift.

Heat extremes

We find a strong increase in the frequency of austral summer months (DJF) warmer than 3-sigma and 5-sigma by the end of the century (2071–2099) (Fig. 1 bottom panels and Fig. ESM.3–4, see Schellnhuber et al. 2014: Appendix A.1

Fig. 1 Temperature changes in LAC for RCP8.5 (4 °C world) for the austral summer months (DJF). Multi-model mean temperature anomalies in degree Celsius (*top row left*) are averaged over the time period 2071–2099 relative to 1951–1980 and normalized by the local standard deviation (*top row right*). Multi-model mean of the percentage of austral summer months (DJF) in the time period 2071–2099 with temperatures greater than 3-sigma (*bottom row left*) and 5-sigma (*bottom row right*)



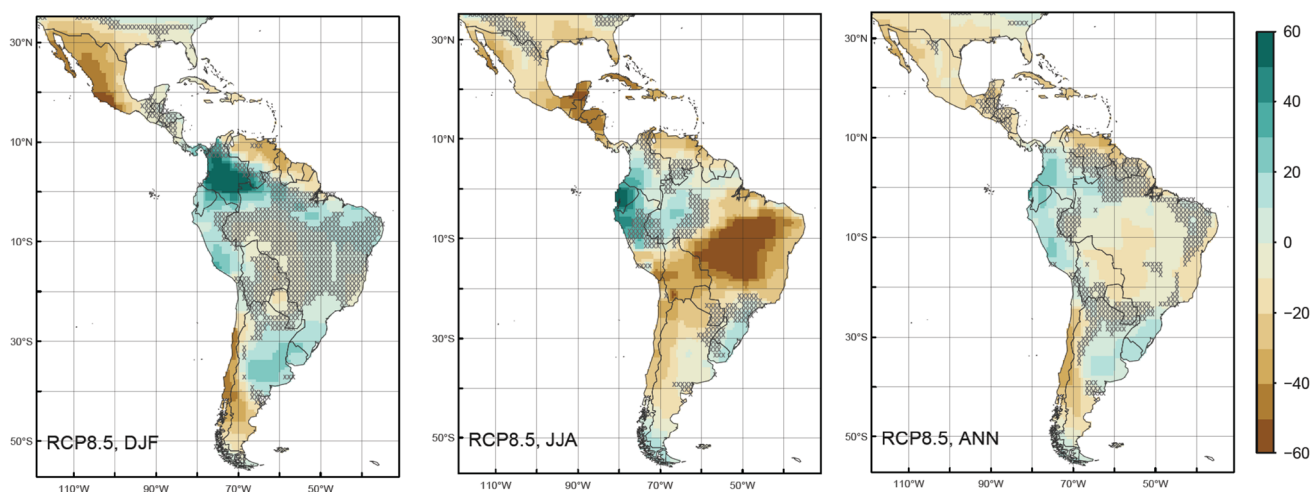


Fig. 2 Multi-model mean of the percentage change in austral summer (DJF, *left*), winter (JJA, *middle*) and annual (*right*) precipitation for RCP8.5 (4 °C world) for Latin America and the Caribbean by 2071–2099 relative to 1951–1980. Hatched areas indicate uncertain results, with two or more out of five models disagreeing on the

for more detailed explanations of 3- and 5-sigma events). The tropics, along the coasts, will see the largest increase in such threshold-exceeding extremes which is confirmed by analyses using the full CMIP5 dataset (Coumou and Robinson 2013; Sillmann et al. 2013a, b). The 5-sigma events, which are absent under present-day climate conditions, will emerge in these countries even in a 2 °C world and are projected to occur in roughly 20 % of summer months. At the same time, 3-sigma events, which are extremely rare today, will be exceeded in roughly half of the summer months during 2071–2099. In a 4 °C world, almost all summer months in the tropics will be warmer than 3-sigma and, in fact, most will be warmer than 5-sigma as well (70 %). Compared to the tropics, the subtropical regions in the north (Mexico) and south (Uruguay, Argentina and southern Chile) are projected to see a more moderate increase in the frequency of threshold-exceeding extremes. In a 2 °C world, 5-sigma events will remain absent and 3-sigma events will still be rare (less than 10 % of summer months). In a 4 °C world, at least half of all summer months are expected to be warmer than 3-sigma by 2071–2099 and 5-sigma events will occur typically in about 20 % of summer months over subtropical regions. This does not mean that absolute warming is less in the subtropical regions or that risks by heat extremes are lower there, but solely that there is greater natural variability in these regions.

Precipitation projections

In a 2 °C world, projected changes in annual and seasonal precipitation for 2071–2099 relative to 1951–1980 are relatively small (± 10 %) and models exhibit substantial disagreement on the direction of change over most land

direction of change. Note that projections are given as percentage changes compared to the 1951–1980 climatology, and thus, especially over dry regions, large relative changes do not necessarily reflect large absolute changes

regions (Fig. ESM.5). However, regions between 10°S and 30°S may experience dry season (JJA) precipitation reductions exceeding 20 % and exceeding 50 % in Mato Grosso, central Brazil (Fig. ESM.5). In a 4 °C world, the models converge in their projections over most regions, but intermodel uncertainty remains over some areas (Fig. 2). Well-defined patterns of change in annual precipitation can be extracted for subregions. Tropical countries on the Pacific coast (Peru, Ecuador and Colombia) are projected to see an increase in annual mean precipitation of about 30 %, with changes in seasonal precipitation exceeding 50 %. Also the Río de la Plata basin is projected to experience a wetting trend. Substantial drying exceeding 10–20 % annual precipitation reductions is projected for Patagonia, the Caribbean and Mesoamerica as well as central Brazil. Notably, there are substantial seasonal differences with JJA precipitation reductions exceeding 40–50 % for Mesoamerica and the Caribbean. The annual mean precipitation in central Brazil is projected to drop by 20 % in a 4 °C world by the end of the century due to a strong and robust decrease in dry season (JJA) precipitation (–50 %). These projected changes in annual and seasonal temperatures generally agree well with those provided by the full set of CMIP5 climate models (Collins et al. 2013) with the notable difference of the central Brazil drying that is less pronounced in the multi-model average of the full CMIP5 set.

Extreme precipitation and droughts

Observations since the 1950s indicate a robust increase in overall precipitation and in intensity of extreme

precipitation events for South America (Skansi et al. 2013). In a 4 °C world, Kharin et al. (2013) find extreme precipitation events (annual maximum daily precipitation with 20-year return values) to intensify by about 25 % over LAC with a large uncertainty range over an ensemble of CMIP5 models. The return time of a 20-year extreme precipitation event in 1985–2005 would reduce to about 6 years (Kharin et al. 2013). These increases are not homogeneous over the full continent. While little-to-not statistically significant, an increase in frequency is projected for the Caribbean, Mesoamerica, southern Argentina and Chile, and hotspots with extreme precipitation increases of more than 30 % are projected in the Serra do Espinhaço in Brazil, the Pampas region in Argentina and the Pacific coastline of Ecuador, Peru and Colombia (Kharin et al. 2013). The latter may be related to an increase in frequency of future extreme El Niño events (Power et al. 2013; Cai et al. 2014). These regions are also found to show the strongest rise in compound maximum 5-day precipitation (which is relevant for flooding events) in a 4 °C world (Sillmann et al. 2013b). Increases in extreme precipitation in southern Brazil and northern Argentina are in line with results from regional climate models (Marengo et al. 2009) and might be dominated by intensification of the South American monsoon system (Jones and Carvalho 2013).

Dai (2012) finds a statistically significant increase in drought conditions for Central America and the Caribbean for the 1950–2010 period, although the significance of this trend depends on the reference period and the formulation of the underlying drought index (Trenberth et al. 2014). Fu et al. (2013) report a significant increase in the length of the dry season over southern Amazonia since 1979. An increase and intensification in meteorological droughts is projected for large parts of South and Central America in a 4 °C world (Sillmann et al. 2013b), although large model uncertainties remain in particular for Central America (Orlowsky and Seneviratne 2013). Accounting for the effects of runoff and evaporation as well as local soil and vegetation, Dai (2012) found that the Amazon basin, Brazil except the southern coast, southern Chile, Central America and northern Mexico are facing severe to extreme drought conditions relative to the present climate by the end of the twenty-first century under the RCP4.5. These results are confirmed by a multi-model analysis in a 4 °C scenario that additionally reveals a strong increase in drought risk in the Caribbean, although uncertainties remain substantial (Prudhomme et al. 2013). Apart from a reduction in precipitation, warming can also cause more arid conditions as enhanced surface temperatures trigger more evapotranspiration—thereby drying the soil (see ESM Text.1 and Fig. ESM.6–7).

Tropical cyclones

Tropical cyclones have important impacts in LAC such as an average 0.83 % drop in economic output after tropical cyclone strikes with large variations between countries (Strobl 2012). Tropical cyclone frequency has increased in the North Atlantic sharply over the past 20–30 years, but uncertainty is large over longer time periods (Bindoff et al. 2013). Kossin et al. (2013) showed a strong and statistically significant increase in lifetime maximum intensity of tropical cyclones over the North Atlantic of 8 m s^{-1} per decade, over 1979–2010, particularly for mid- to high-intensity storms. Such observed changes were shown to be linked to both anthropogenic climate change and internal climate variability (Camargo et al. 2012; Villarini and Vecchi 2013; Wang and Wu 2013). Differential warming of the tropical Atlantic, with historically observed warming higher than average for the tropics, tends to enhance tropical cyclone intensification in the region (Knutson et al. 2013). No significant trends have been observed over the eastern North Pacific (Kossin et al. 2013), but in general tropical cyclones have been observed to migrate polewards (Kossin et al. 2014).

Projections of tropical cyclone frequency and intensity are difficult, because the interplay of several factors is unclear (Bindoff et al. 2013). El Niño events tend to enhance wind shear over the Gulf of Mexico and the Caribbean Sea and thus suppress Atlantic tropical cyclones (Arndt et al. 2010; Aiyer and Thorncroft 2011; Kim et al. 2011). On the other hand, El Niño events have been shown to increase tropical cyclone activity in the eastern North Pacific (Kim et al. 2011; Martinez-Sanchez and Cavazos 2014). Observational evidence, however, suggests atmospheric patterns tend to steer tropical cyclones away from the Mexican coast during El Niño years (and toward the coast in La Niña years), so that the net effect on the Pacific coastlines of the Americas remains unclear. In addition to such dynamic changes, thermodynamic processes alone can also work to suppress tropical cyclone formation and intensification (Mallard et al. 2013).

In the long term, simulations from a range of models show that tropical cyclone frequency will not be affected much by continued global warming but mean intensity, as well as the frequency of the most intense tropical cyclones, is projected to increase (Knutson et al. 2010; Tory et al. 2013; Stocker et al. 2013). Using CMIP5 models (50 % uncertainty range across 17 GCMs), Villarini and Vecchi (2013) projected that the Power Dissipation Index would increase by 100–150 % in a 2 °C world over the North Atlantic. A considerably larger increase and a much wider range of about 125–275 % were projected for a 4 °C world.

Bender et al. (2010) used a variety of models to initialize a very high-resolution operational hurricane prediction model, noting an increase of 80 % in the frequency of the strongest category 4 and 5 Atlantic tropical cyclones for about 3.4 °C warming (a 3 °C world). Knutson et al. (2013) also found an 80 % increase in the strongest category tropical cyclones for the same warming level and class of models and around a 40 % increase for roughly 1.5–2.5 °C warming (early and late twenty-first century, RCP4.5). The eastern North Pacific is less well represented in the scientific literature and studies projected either no significant trends for this region under future climate change (Murakami et al. 2011, 2012) or an increase in frequency of tropical cyclones (particularly large near the coast of southeast Mexico) (Emanuel 2013). With projected increased intensity and frequency of the most intense storms, and increased atmospheric moisture content, an increase of 10 % in the rainfall intensity averaged over a 200-km radius from the tropical cyclone center for the Atlantic, and an increase of 20–30 % for the tropical cyclone's inner core by the end of the twenty-first century for roughly 2.5–3.5 °C global warming is estimated (Knutson et al. 2010, 2013).

Regional sea-level rise

In LAC, global mean sea-level rise dominates the regional sea-level signal (Fig. 3, Fig ESM.8). Still, regional variation exists with generally higher projected sea-level rise at the Atlantic than at the Pacific coast and an increasing ice-

sheet contribution toward lower latitudes. Due to a robust southeasterly trade wind intensification over the Southern Pacific and associated cold water upwelling (Timmermann et al. 2010; Merrifield and Maltrud 2011), sea-level rise below the global mean is projected for the Southern Pacific coast (median estimate for Valparaiso: 0.55 m for a 4 °C world cf. Table ESM.4). In contrast, Recife on the Atlantic coast of Brazil is projected to experience above-average sea-level rise (median estimate 0.63 m). The sea-level rise at the continental Caribbean coast exceeds the projection for the Caribbean islands (Barranquilla, median estimate 0.65 for a 4 °C world). The difference may be linked to a weakening of the Caribbean Current that is connected to the Atlantic meridional overturning circulation (Pardaens et al. 2011).

Regional impacts

Glacial retreat and snowpack changes

Andean glaciers are shrinking because of increased melt rates, decreased accumulation, changes in the ice dynamics and/or a combination of all these factors (Lopez et al. 2010; Ivins et al. 2011; Jacob et al. 2012; Marzeion et al. 2012; Giesen and Oerlemans 2013; Schaefer et al. 2013; Table ESM.1). Across several studies and different methods, a clear change in glacier evolution can be seen after the late 1970s, accelerating in the mid-1990s and again in the early 2000s (Rabatel et al. 2013). This acceleration

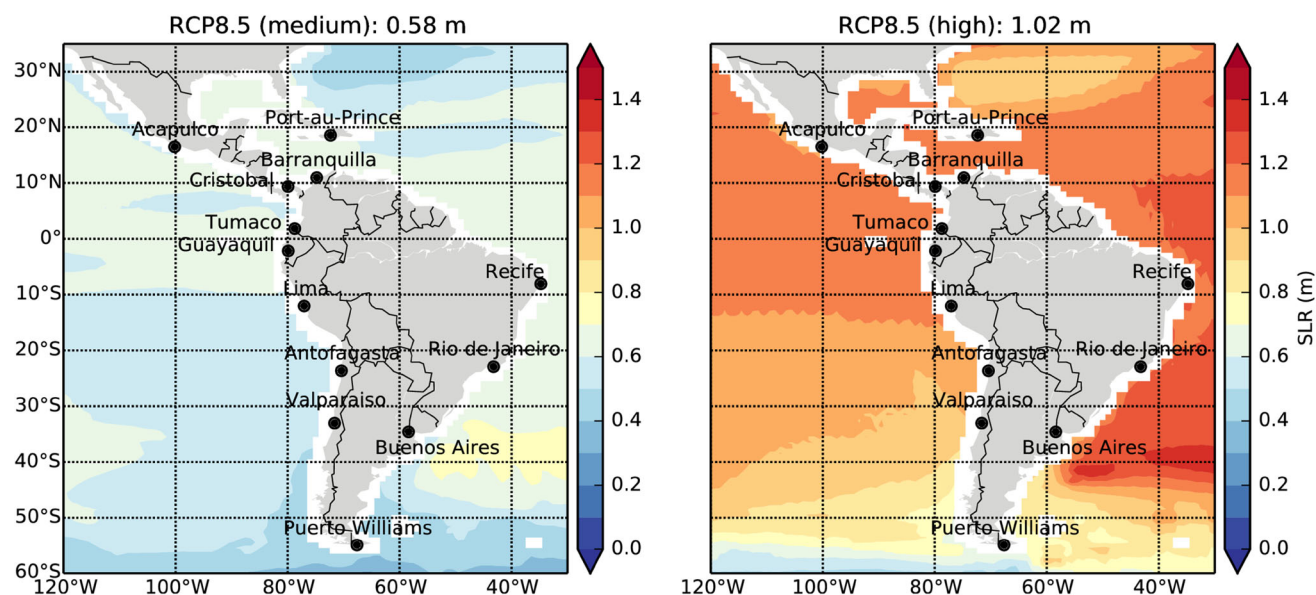


Fig. 3 Patterns of regional sea-level rise. Median (*left column*) and high (*right column*) estimates of projected regional sea-level rise for the RCP8.5 scenario (4 °C world) for the period 2081–2100 relative

to the reference period 1986–2005. Associated global mean rise is indicated in the panel titles. Representative cities are denoted by *black dots* and *numbers* provided in Table ESM.4

started slightly delayed in the 1990s for glaciers located at mid- or high latitudes.

In the subtropical Andes of Chile and western Argentina, there is no significant trend in snowfall over 1951–2004 (Masiokas et al. 2006, 2012). However, the data display a marked inter-annual variability ranging from 6 to 257 % around the 1966–2004 mean, with a clear influence from the warm phases of ENSO. Studies on snowpack in the southern Andes are rare, but changes in snowpack extent magnify changes in the seasonality of the water availability by reducing dry season flows and increasing wet season flows (Vicuña et al. 2013).

Tropical glaciers are projected to lose at least 66 % of their volume in a 3 °C world and to disappear almost completely in a 4 °C warmer world (Marzeion et al. 2012; Giesen and Oerlemans 2013; Radić et al. 2013). The much larger glaciers in the southern Andes are less sensitive to global warming and are projected to shrink by at least 21 % in a 2 °C world up to 72 % in a 4 °C world (see Table ESM.1). The models used in these studies rely on a scaling methodology which may overestimate the recession of the small remnant tropical glaciers. The accelerated melting will lead to increasing runoff, but when the glacier reservoirs disappear runoff will also decrease, particularly in the dry season. Following the trend in the tropical Andes (Poveda and Pineda 2010), this peak is expected within the next 50 years (Chevallier et al. 2011) if it has not already occurred (Baraer et al. 2012). Reliable projections for snowpack and snow cover changes in the Andes are lacking.

Water resources, water security and floods

LAC has abundant water resources, but their distribution is temporally and regionally unequal (Magrin et al. 2007). Due to the unreliable rainfall, groundwater resources and water from glacier and snowmelt play a crucial role in supplying local water (Vuille et al. 2008; Chevallier et al. 2011; Hirata and Conicelli 2012). Moreover, LAC suffers from widespread floods and landslides (Maynard-Ford et al. 2008). Heavy precipitation events in the context of ENSO or tropical cyclones can lead to disastrous floods, especially in regions with steep terrain (Mata et al. 2001; Poveda et al. 2001; Mimura et al. 2007; IPCC 2012). Coastal areas in the Caribbean and Central America suffer from flooding as a result of storm surges and tropical cyclones (Dilley et al. 2005; Woodruff et al. 2013). In the Andes, glacial lake outbursts present a permanent hazard for Andean cities (Chevallier et al. 2011; Carey et al. 2012).

In Central America, there is a high agreement on decreasing mean annual runoff and discharge, although the magnitude of the change varies (Milly et al. 2005; Maurer et al. 2009; Imbach et al. 2012; Arnell and Gosling 2013; Hidalgo et al. 2013; Nakaegawa et al. 2013; Schewe et al.

2013). The trend seems to be more pronounced for the northern than for the southern part of Central America (Imbach et al. 2012; Hidalgo et al. 2013). The Caribbean lacks long-term measured streamflow data, and runoff projections are therefore of low confidence (FAO 2003; Hidalgo et al. 2013). However, freshwater availability may decrease due to a combination of lower precipitation, high abstraction rates and sea-level rise leading to an intrusion of sea water into coastal aquifers (Mimura et al. 2007; Cashman et al. 2010). Although floods often seem to be associated with land-use change, more severe flooding events may also occur with climate change, e.g., related to tropical cyclones (Cashman et al. 2010; IPCC 2012). Higher discharge seasonality is projected for the tropical Andes. Lower dry season discharge has already been observed during the past two decades (Baraer et al. 2012). Streamflows during the dry season are projected to decrease strongly because of ongoing glacier retreat and snowmelt decrease (Juen et al. 2007; Baraer et al. 2012; Kinouchi et al. 2013). However, streamflow during the wet season may increase due to an increase in direct runoff from non-glaciered areas (Juen et al. 2007; Kinouchi et al. 2013). The region has a high flood and landslide risk which is projected to increase (Carey 2005; Hirabayashi et al. 2013). For the Central Andes, a trend toward an earlier snowmelt season and timing of the center of mass of flows was observed and projected locally (Cortés et al. 2011; Vicuña et al. 2013; Demaria et al. 2013). Lower dry season discharges may cause significant water supply problems in downstream and urban areas (Vuille et al. 2008; Viviroli et al. 2011; Masiokas et al. 2013), with poorer areas most affected (Buytaert and De Bièvre 2012), and might endanger electrical power generation (Seoane and López 2007). Runoff projections for the Amazon are uncertain due to the high variability of rainfall projections using different GCMs and uncertainties introduced by hydrological impact models (Buytaert et al. 2009; Exbrayat et al. 2014). Guimberteau et al. (2013) found that, especially in the south, low flows become more pronounced by the middle of this century with 2 °C global warming. Nakaegawa et al. (2013) found total annual runoff decreases in the southern half of the Amazon River in a 3 °C world. However, for the western part of the basin, a likely increase in streamflow, runoff, flood zone and inundation time was projected (Guimberteau et al. 2013; Langerwisch et al. 2013; Mora et al. 2013b). The direction of discharge and groundwater recharge trends in northeast Brazil vary due to diverging rainfall projections (Krol and Bronstert 2007; Montenegro and Ragab 2010; Döll and Schmied 2012; Portmann et al. 2013; Schewe et al. 2013). The Río de la Plata region experienced a 10–30 % increase in river runoff during the twentieth century (García and Vargas 1998; Jaime and Menéndez 2002; Menéndez and Berbery 2005;

Milly et al. 2005). There are no consistent river runoff projections for the basin because the directions of rainfall projections vary (Milly et al. 2005; Nóbrega et al. 2011; Bravo et al. 2013; Nakaegawa et al. 2013). Camilloni et al. (2013) projected an increase in frequency and duration of river flooding in a >3 °C world in the Uruguay and Paraná basins. Hirabayashi et al. (2013) showed a decrease in the twentieth century 100-year return period for floods for the Parana in a 4 °C world, but there was little consistency across the 11 GCMs used. A decrease in mean runoff is projected for southernmost South America (Milly et al. 2005; Schewe et al. 2013).

Agricultural yields

Climate change impacts on crop yields vary depending on crop type and location, but most projections show negative climate change impacts (summarized in Table ESM.1). Yield declines are projected for wheat, soybeans and maize in several countries (Fernandes et al. 2012). Coffee farming might have to migrate to higher altitudes or other cultivation regions to maintain present yields (Camargo 2010; Laderach et al. 2011; Zullo et al. 2011). Yield projections for rice and sugarcane mostly increase (Fernandes et al. 2012; Marin et al. 2012). In several studies, including CO₂ fertilization increases yields or only leads to small negative yield changes for C₃ plants but not for C₄ plants (Costa et al. 2009; ECLAC 2010; Lapola et al. 2011; Rosenzweig et al. 2013).

A meta-analysis of the impacts of climate change on crop yields for LAC (see Schellnhuber et al. 2014; Appendix A.3) reveals no significant influence of temperature increase on crop yields across all available studies but a significant positive relationship between crop yield change and temperature when CO₂ fertilization is considered (see Fig. 4, Table ESM.5). However, the beneficial effects of CO₂ fertilization are uncertain (e.g., Ainsworth et al. 2008). If the effects of CO₂ fertilization are not considered, the relationship remains significant but becomes negative, with increasing temperature leading to considerable yield declines (see Fig. 4, Table ESM.5).

The impact of ozone on crop yields has been neglected in many climate impact projections although it could offset the benefits of CO₂ fertilization in C₃ plants (and even lead to a yield reduction in C₄ plants) (Jaggard et al. 2010). By 2030, increasing surface ozone could decrease yields in Latin America by up to 7.8 % for wheat, 2.9 % for maize and 7.5 % for soybeans depending on the emissions levels of ozone precursors (Avnery et al. 2011). Moreover, the effects of plant diseases are uncertain but potentially negative (Ghini et al. 2011; Luck et al. 2011; Porter et al. 2014). For example, Coffee leaf rust (*Hemileia vastratix*) and soybean rust (*Phakopsora pachyrhizi*) are expected to move further south and affect South American countries (Alves et al.

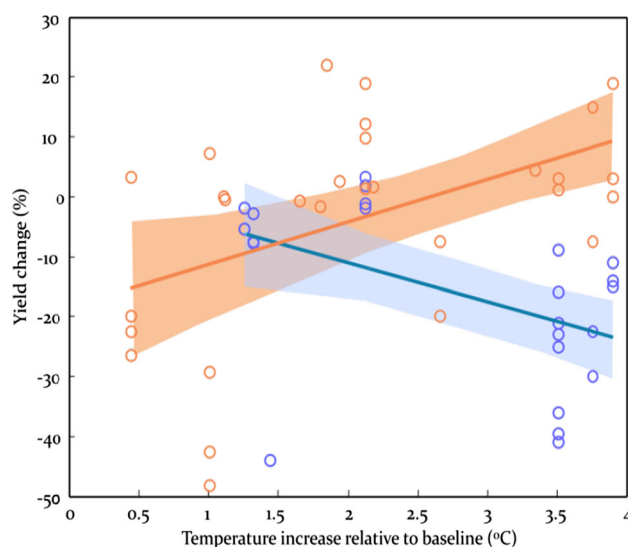


Fig. 4 Meta-analysis of crop yield reductions. Best-fit line for LAC studies not considering the effects of adaptation measures or those of CO₂ fertilization (*blue line*) and for studies considering the effects of CO₂ fertilization (but no adaptation, *orange*) and their 95 % confidence intervals of regressions consistent with the data based on 500 bootstrap samples (*patches*) (color figure online)

2011). Estay et al. (2009) projected an increase in insect population densities of grain pests in Chile of 10–14 % in a 3 °C world and 12–22 % in a 4 °C world.

Livestock

The livestock sector in LAC is of high economic importance, especially in Brazil and Argentina (ECLAC et al. 2012), but climate impact studies are scarce. Climate change can impact the quantity and quality of livestock feed, and heat stress directly affects livestock productivity. Heat stress is known to reduce cattle food intake and milk production and also to affect reproduction, growth and mortality rates (Porter et al. 2014). In a 1–2 °C warmer world, livestock species choice (i.e., the adoption of new livestock) is projected to mostly decline or marginally increase across Argentina, Brazil, Chile, Colombia, Ecuador, Uruguay and Venezuela for beef cattle, dairy cattle, chicken and pigs (Seo et al. 2010). Meanwhile, the adoption of sheep is projected to increase by up to 20 % because sheep are better adapted to warmer and drier conditions (Seo et al. 2010). In Paraguay, beef cattle production is projected to decrease by 16–27 % in scenarios leading to 2 and 3 °C warming (ECLAC 2010).

Biodiversity

South America is a biodiversity hotspot (MEA 2005; Myers et al. 2000). Habitat destruction and fragmentation

by land-use change as well as the commercial exploitation of species groups are currently larger threats to biodiversity than climate change (e.g., Hof et al. 2011), but climate change is projected to become increasingly important for species loss (MEA 2005; Vuuren et al. 2006). Future species loss is difficult to quantify because most biodiversity models do not take biotic interactions (e.g., food-web interactions, species competition) and resource limitations into account.

The G200 ecoregions (Olson and Dinerstein 2002) located in LAC may experience severe climate change in the future (Beaumont et al. 2010; Li et al. 2013). Further, 38.4 and 11.5 % of the surface of the biodiversity hotspot of Tumbes-Choco-Magdalena and the Mesoamerican biodiversity hotspot, respectively, will be experiencing non-analog climates in a warmer than 2 °C world (García-López and Allué 2013). Heyder et al. (2011) project a range of small to severe ecosystem changes for the whole South American continent in a 2 °C and warmer world. In a 4 °C world, results of one dynamic vegetation model show severe ecosystem changes for more than 33 % of the area in 21 out of 26 distinct biogeographic regions in South America (Gerten et al. 2013). Warszawski et al. (2013) projected such severe ecosystem changes in a 3 °C world in South America (notably in Amazon, Guyana moist forests and Brazilian Cerrado) when applying an ensemble of seven dynamic vegetation models. Bellard et al. (2014) projected that out of 723 Caribbean islands, 63 and 356 will be entirely submerged under one and six meters of sea-level rise, respectively. A one-meter sea-level rise is within the range of sea-level rise projected in a 4 °C world (see section “Regional sea-level rise”).

Little is known about the consequences of future climate change on specific taxa in the region. There are some studies on insects (Deutsch et al. 2008), amphibians (Lawler et al. 2009; Sinervo et al. 2010; Hof et al. 2011; Loyola et al. 2013; Mesquita et al. 2013; Lemes et al. 2014), sea turtles (Fish et al. 2005), birds (Anciães and Peterson 2006; Jetz et al. 2007; Marini et al. 2009; Souza et al. 2011), marsupials (Loyola et al. 2012), mammals (Schloss et al. 2012; Torres et al. 2012) and plant species (Feeley et al. 2012; Simon et al. 2013). Overall, most studies project range contractions or inability to keep pace with climate. Across different species, Thomas et al. (2004) project increasing extinction rates with increasing warming levels.

Amazon rainforest degradation and tipping point

Old-growth rainforests in the Amazon store approximately 100 billion tons of carbon in their biomass (Malhi et al. 2006; Saatchi et al. 2011) and recycle precipitation through evapotranspiration, thus contributing to local rainfall

(Zemp et al. 2014). A loss of these forests due to climate change and deforestation would release an enormous amount of carbon into the atmosphere and reduce their evapotranspiration potential leading to strong climate feedbacks (Betts et al. 2004; Cox et al. 2004; Costa and Pires 2010) and a potential tipping point (Lenton et al. 2008). A critical tipping point has been identified at around 40 % deforestation, when altered water and energy feedbacks between remaining tropical forest and climate may lead to a decrease in precipitation (Sampaio et al. 2007). Recent evidence from a large-scale and long-term experiment suggests that the feedbacks between climatic extreme events such as droughts and forest fires increase the likelihood of an Amazon dieback (Brando et al. 2014). However, the most recent modeling studies suggest that the Amazon dieback is an unlikely, but possible, future for the Amazon region (Good et al. 2013).

The 2005 and 2010 Amazon droughts lead to increased tree mortality and reduced tree growth due to water stress (Lewis et al. 2011) possibly reversing the role of the intact forest as a carbon sink (Phillips et al. 2009; Lewis et al. 2011). Two multi-year rainfall exclusion experiments in Caxiuanã and Tapajós National Forest demonstrated that once deep soil water is depleted, wood production is reduced by up to 62 %, aboveground net primary productivity declines by 41 %, and mortality rates for trees almost double (Nepstad et al. 2007; Brando et al. 2008; Costa and Pires 2010). Thus, an increase in extreme droughts or a prolonged dry season (Fu et al. 2013) may have the potential to cause large-scale forest dieback.

Deforestation and forest degradation are also factors which crucially influence future changes in vegetation carbon. Gumpenberger et al. (2010) found relative changes in carbon stocks of –35 to +40 % in a protection scenario without deforestation and –55 to –5 % with 50 % deforestation in a 4 °C world. Poulter et al. (2010) found a 24.5 % agreement of projections for a decrease in biomass in simulations with 9 GCMs in a 4 °C world.

Studies projecting future fires in the Amazon are scarce. Fires are projected to increase along major roads in the southern and southwestern part of Amazonia with a 1.8 °C global warming by 2040–2050 (Silvestrini et al. 2011; Soares-Filho et al. 2012). High rates of deforestation would contribute to an increasing fire occurrence of 19 % by 2050, whereas climate change alone would account for a 12 % increase (Silvestrini et al. 2011).

Despite a large number of studies (see Table ESM.1 for a summary), the identification of the processes and the quantification of thresholds at which a tipping point is triggered (e.g., a potential transition from forest to savannah) are still incomplete. Recent analyses have downgraded the probability from 21 to 0.24 % for a 4 °C regional warming when coupled carbon-cycle climate

models are adjusted to better represent the inter-annual variability of tropical temperatures and related CO₂ emissions (Cox et al. 2013). This holds true, however, only when the CO₂ fertilization effect is realized as implemented in current vegetation models (Rammig et al. 2010). Moreover, large-scale forest degradation as a result of increasing drought may already impair ecosystem services and functions without a forest dieback necessarily occurring and will impact forest dwelling/depending communities.

Fisheries

Anthropogenic changes in temperature, salinity, oxygen content and pH levels have been observed for oceans over the past 60 years (Pörtner et al. 2014). In LAC, small-scale fishers and people dependent on large-scale fishing and associated industries may be particularly affected by climate change if they are unable to diversify their livelihoods. Peru and Colombia are among the countries whose fisheries are most vulnerable to climate change (Allison et al. 2009; Magrin et al. 2014).

In response to changing oceanic conditions, fish stocks have been observed in, and are further expected to shift to, higher latitudes (Perry et al. 2005). A global study that considers the habitat preference of 1066 commercially caught species and projects changes to primary productivity, computes the expected changes in fish species distribution and regional patterns of maximum catch potential by 2055 in a scenario leading to warming of approximately 2 °C in 2050 (and 4 °C by 2100) (Cheung et al. 2010). Their results for LAC indicate a mixed picture: Catch potential is expected to increase up to 100 % further offshore of the southern part of Latin America and to decrease by 15–50 % along the Caribbean coasts, by 5–50 % in Caribbean waters and parts of the Atlantic coast of Central America, by more than 50 % off the Amazonas estuary and the Rio de la Plata and by up to 30 % along the coasts of Peru and Chile (with increases toward the south). These projections are, however, uncertain because expected declines in ocean pH (ocean acidification), direct human pressures and local processes, which escape the coarse resolution of global models, are not taken into account (Cheung et al. 2010). Incorporating the effects of decreasing ocean pH and reduced oxygen availability yields catch potentials that are 20–30 % lower relative to simulations not considering these factors in the northeast Atlantic (Cheung et al. 2011). Considering the effects of species interaction on redistribution and abundance, Fernandes et al. (2013) report latitudinal shifts in the North Atlantic to be 20 % lower than reported by the bioclimatic envelope model developed by Cheung et al. (2010). For the Humboldt Current System, Blanchard et al. (2012)

projected a 35 % decline in phytoplankton and zooplankton density and similar magnitudes of change in the overall biomass of fish under 2 °C global warming by 2050.

Coral reefs

Coral reefs provide ecosystem services which are locally important for subsistence fisheries, income from tourism and protection from coastal storm surges (Hoegh-Guldberg et al. 2007). Coral bleaching events on a large scale have been linked to unusually high sea-surface temperatures but also pollution, overfishing and the related shift in species composition are important factors (De'ath et al. 2012). Hurricanes have been found to cool waters, alleviating heat stress and thereby reducing the risk of bleaching (Eakin et al. 2010). This effect could temporarily outweigh the negative effects of direct damage (e.g., through breakage) (Carrigan and Puotinen 2014 but see Gardner et al. 2005).

Due to decreasing availability of calcium carbonate and increasing sea-surface temperatures, Meissner et al. (2012) projected that most coral reef locations in the Caribbean Sea and western Atlantic will be subject to a 60–80 % probability of annual bleaching events with 2 °C warming by 2050, with areas at the coast of Guyana, Suriname and French Guiana being exposed to a 100 % probability. In contrast, under 1.5 °C warming by 2050, most locations in the Caribbean Sea have a comparably low risk of 20–40 % probability of annual bleaching events, with the waters of Guyana, Suriname, French Guiana and the north Pacific being at slightly higher risk (up to 60 % probability). By the year 2100, almost all coral reef locations are expected to be subject to severe bleaching events occurring on an annual basis in a 4 °C world. Exceptions are major upwelling regions, which experience a risk of 50 %. With warming leading to a 2 or 4 °C world, the median year in which bleaching events start to occur annually is 2046 or 2040, respectively (Van Hooijdonk et al. 2013). Generally, the reefs in the northern waters of the Caribbean Sea appear to be less sensitive than those in the south. However, reefs at the higher latitude fringes of the tropical coral range (both north and south) are likely to be more heavily affected by ocean acidification (Caldeira 2013). Different scenarios leading to 2, 3 and 4 °C worlds showed little difference in coral cover (Buddemeier et al. 2011): By 2020, live coral reef cover is projected to have halved from its initial state; by the year 2050, live coral cover is less than 5 %; in 2100, it is less than 3 %. Assuming coral adaptation by gaining an additional 1 °C of heat tolerance, the loss of live coral cover below 5 % is prolonged by around 30 years. A 90 % loss of coral reef cover would lead to direct economic losses of \$8.712 billion (2008 value) (Vergara et al. 2009). While there are limitations to the projections of coral reef future, a bleak picture emerges

from the studies available, which is in accordance with global studies showing that the global mean temperature at which 90 % of coral reefs are at risk of extinction is 1.5 °C above pre-industrial levels (Frieler et al. 2012).

Human health

Among the main human health risks in LAC are vector-borne diseases such as malaria, dengue fever, leishmaniasis and fascioliasis, and food- and water-borne diseases such as cholera and childhood diarrheal disease. Climate change is expected to play a contributing role in determining the incidence of dengue fever (Confalonieri et al. 2007), although it is difficult to separate the impact of climatic factors from that of urbanization and population mobility (Barclay 2008). Studies from Mexico (Hurtado-Diaz et al. 2007) and Puerto Rico (Johansson et al. 2009) show a correlation between increases in rainfall and temperature and increased incidence of the disease. In Brazil in the period 2001–2009, a 1 °C increase in monthly minimum temperature was associated with a 45 % increase in dengue fever cases the following month and a 10 mm increase in precipitation with a 6 % increase (Gomes et al. 2012). Projections by Colon-Gonzalez et al. (2013) point to an upsurge in dengue incidence in Mexico of 18 % by 2030, 31 % by 2050 and 40 % by 2080 with a warming scenario leading to a 4 °C world by 2100.

Evidence shows an increasing spread of malaria to higher elevations in northwest Colombia during the last three decades due to rising temperatures (Siraj et al. 2014), and it is possible the disease will spread into other high-altitude areas such as the cities of Quito and Mexico City (Moreno 2006). The connection between malaria and climate change is unclear, however, due to the complexity of the factors involved (e.g., land use, domestic water storage patterns, vector control programmes). It is likely that the effect of climate change on malaria patterns will not be uniform, with increased incidence in some areas but declines in places where decreases in precipitation are projected, including parts of the Amazon and Central America (Haines et al. 2006). Caminade et al. (2014) projected a lengthening of the malaria transmission season in the highlands of Central America and southern Brazil by the 2080 s, but a shortening in the tropical regions of South America.

Climatic variables have been shown to be decisive in determining the extent of cholera outbreaks (Koelle 2009). A recent study in Haiti shows that increased rainfall is followed by increased cholera risk 4–7 days later (Eisenberg et al. 2013). The relative risk of diarrheal disease in South America is expected to increase by 5–13 and 14–36 % for the period 2010–2039 and 2070–2099 with 1.3 and 3.1 °C warming, respectively (Kolstad and Johansson 2011).

Coastal infrastructure

Hallegatte et al. (2013) found that, by 2050, coastal flooding could generate approximately \$940 million of mean annual losses in the 22 largest coastal cities in LAC with a sea-level rise of 20 cm and about \$1.2 billion with a sea-level rise of 40 cm. The damage to coastal infrastructure associated with tropical cyclones making landfall is also projected to change (Hallegatte 2007; Mendelsohn et al. 2011). In a scenario leading to a 4 °C world and featuring a 0.89- to 1.4-m sea-level rise, tropical cyclones in the Caribbean alone could generate an extra \$22 billion and \$46 billion in storm and infrastructure damages and tourism losses by 2050 and 2100, respectively, compared to a scenario leading to a 2 °C world (Bueno et al. 2008). Cumulative losses induced by increasing tropical cyclone intensity of 2 and 5 % compared to average values from 1995 to 2006 could increase to about \$110 and \$114 billion, respectively, during the period 2020–2025 in the Caribbean, Central America and Mexico (Curry et al. 2009). The potential increase in tropical cyclone intensity may increase ships' port downtime and therefore increase shipping costs (Esteban et al. 2012; Chhetri et al. 2013). Impacts on seaports will also have indirect consequences on local economies as import disruptions generate price increases for imported goods and export disruptions decrease revenues and incomes (Becker et al. 2012). Beach tourism is particularly exposed to several climate change stressors, including sea-level rise, modified tropical storm pattern and heightened storm surges (Simpson et al. 2011). In Jamaica, for example, coastal tourist resorts are two-to-three times more exposed to climate change-related stressors than inland touristic resorts (Hyman 2013).

Energy systems

LAC countries have a diverse energy mix (see Table ESM.6), but with projected changes in water availability (cf. “Glacial retreat and snowpack changes” and “Water resources, water security, and floods” sections), thermal electricity plant cooling systems may become less efficient (Mika 2013; Sieber 2013) and also hydroelectric power generation will be affected (Hamududu and Killingtveit 2012). In Peru, it is estimated that a 50 % reduction in glacier runoff would result in a decrease in annual power output of approximately 10 %, from 1540 to 1250 GWh (Vergara et al. 2007). Hamududu and Killingtveit (2012) found that production will increase by 0.30 TWh (or 0.03 %) in the Caribbean compared to 2005 production levels, and by 0.63 TWh (or 0.05 %) in South America, under 2 °C global warming by the middle of the twenty-first century. An increase in frequency of low flows in scenarios leading to a 2 °C and 3 °C world implies a

proportional decrease in hydropower capacity for the two main large reservoirs used for hydroelectricity generation in El Salvador, reducing the economic return from the existing facility and threatening the return on investments in future hydroelectric infrastructures (Maurer et al. 2009). For Brazil, de Lucena et al. (2009) project that average annual river flows will decrease by 10.8 % with 2.9 °C global warming, and by 8.6 % with 3.5 °C global warming, by 2071–2100. For the Rio Grande River basin, the difference between the lowest (−20 %) and the highest (+18 %) estimates of average river flow with a global warming of 2.1 °C depending on the GCM chosen highlights the limitations of current models to project the potential hydropower production (Nóbrega et al. 2011). Popescu et al. (2014) showed an increase in the maximum hydropower energy potential for the La Plata basin of between 1 and 26 % with a global warming of 1.8 °C by 2031–2050.

The results of these studies need to be interpreted with care. Global studies such as from Hamududu and Killington (2012) do not take into account seasonality and

impacts of climate change on the timing of river flows, potential spatial variability and changes occurring over short distances, potential impacts of floods and droughts or impacts on river runoff from decreasing snow cover and snowfall. Such differences may explain, at least partly, the differences between the significant decrease in hydropower capacity at the micro-level as projected by Maurer et al. (2009) and the increase in hydropower generation at the macro-level projected by Hamududu and Killington (2012). Similarly, De Lucena et al. (2009) also only accounted for the average behavior of flows and did not integrate potential change in seasonality or the effects of extreme dry or wet events on hydropower generation.

Implications for regional development

This paper shows that the Latin America and Caribbean region will be severely affected by climate change, even under lower levels of warming (Fig. 5, Table ESM.1). The biophysical impacts described here interact with the

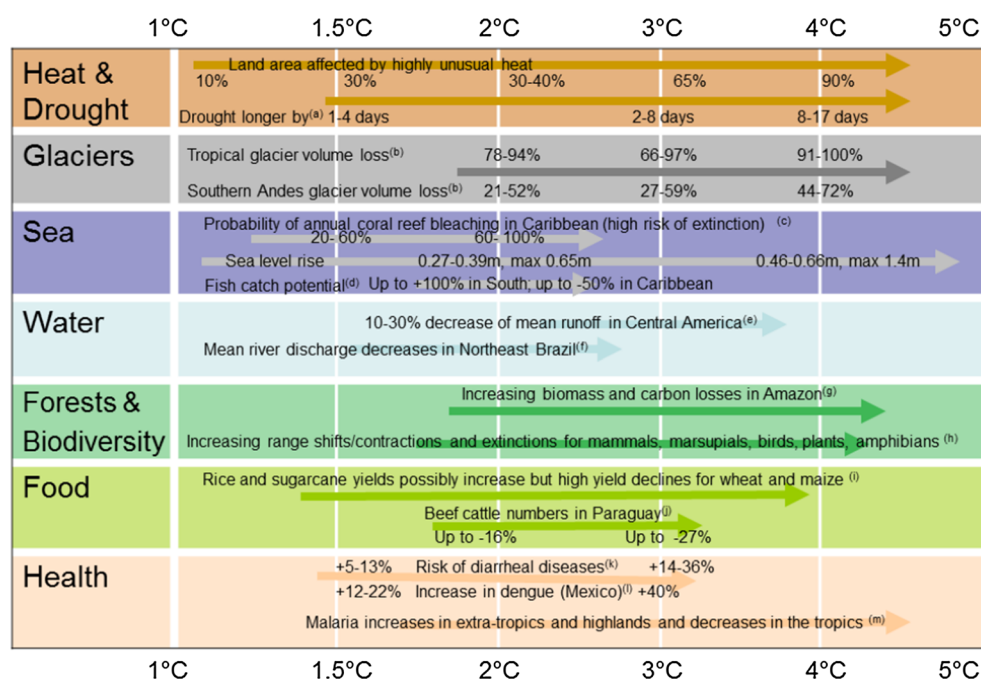


Fig. 5 Projected impacts of climate change in key sectors in the Latin America and Caribbean region. Warming levels are relative to pre-industrial temperatures. The impacts shown here are a subset of those summarized in this paper and in Table ESM.1. The arrows indicate solely the range of warming levels assessed in the underlying studies, but do not imply any graduation of risk unless noted explicitly. In addition, observed impacts or impacts occurring at lower or higher levels of warming that are not covered by the key studies highlighted here are not presented (e.g., coral bleaching already occurs earlier than 1.5 °C warming, but the studies presented here only start at 1.5 °C). Adaptation measures are not assessed here although they can be crucial to alleviate impacts of climate change.

The layout of the figure is adapted from Parry (2010). The lower-case superscript letters indicate the relevant references for each impact. If there is no letter, the results are based on additional analyses for this report. (a) Sillmann et al. (2013b); (b) Marzeion et al. (2012); Giesen and Oerlemans (2013); Radic et al. (2013); (c) Meissner et al. (2012); (d) Cheung et al. (2010); (e) Hidalgo et al. (2013); (f) Döll and Schmied (2012); (g) several studies without considering CO₂ fertilization, see Table ESM.1; (h) several studies, see Table ESM.1; (i) several studies, see Table ESM.1; (j) ECLAC (2010); (k) Kolstad and Johansson (2011); (l) Colon-Gonzalez et al. (2013); (m) Béguin et al. (2011); Caminade et al. (2014); Van Lieshout et al. (2004)

existing vulnerabilities in the region (“Social, economic and demographic profile of the Latin America and Caribbean region and vulnerabilities to climate change” section) and may affect development and human well-being in LAC in a variety of ways: Changes to the hydrological cycle endanger the stability of freshwater supplies and ecosystem services on which many people in LAC depend. Extreme events will strongly affect the rural and urban poor who often reside in informal settlements in high-risk areas (e.g., flood plains and steep slopes). Intense rainfall events can quickly overwhelm natural drainage channels in the landscape as well as urban drainage systems that are unlikely to have been designed for the possible increased intensity of future rainfall events. At lower levels of warming, glacial melt in the Andes will reduce freshwater and hydropower during the dry season for communities and large Andean cities which are often important economic centers, while increasing the risks of flooding in the short term and impacting agriculture and environmental services downstream. More intense tropical cyclones would interact adversely with rising sea levels, exacerbating coastal flooding and storm surge risks, putting entire economies and livelihoods of island states at risk. The Caribbean is particularly vulnerable as more than 50 % of its population lives along the coast and around 70 % live in coastal cities. Degrading coral reefs will endanger tourism revenues and undermine biodiversity, fisheries and the protection of coastal zones. Climate change could also place at risk small-scale subsistence agriculture and large-scale agricultural production for export.

Finally, our assessment of climate change impacts in different sectors shows that impacts are likely to occur simultaneously and possibly interact, thus increasing the risks for development when, e.g., crop yield is reduced, transport disrupted and houses of workers and production infrastructure damaged. Although, climate change challenges those already vulnerable and disfavored the most, large businesses could also suffer which in turn has implications for a large labor force, threatening population groups currently considered less vulnerable. Studies of interacting and cascading impacts and how these affect different population groups and infrastructure types are needed to gain a clearer picture of the relationship between climate change and development, as well as to better plan and implement mitigation and adaptation activities. A key challenge is to properly link biophysical impacts, often projected until the end of this century, to development issues which are highly dynamic and hard to project at similar timescales as the climate impact projections.

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